

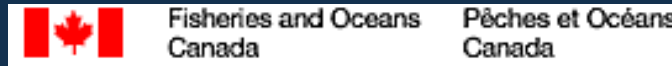
The Arctic Freshwater Balance: A Network Perspective

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AWI



Freshwater and Arctic Change

A Framework for Quantifying & Understanding Change:

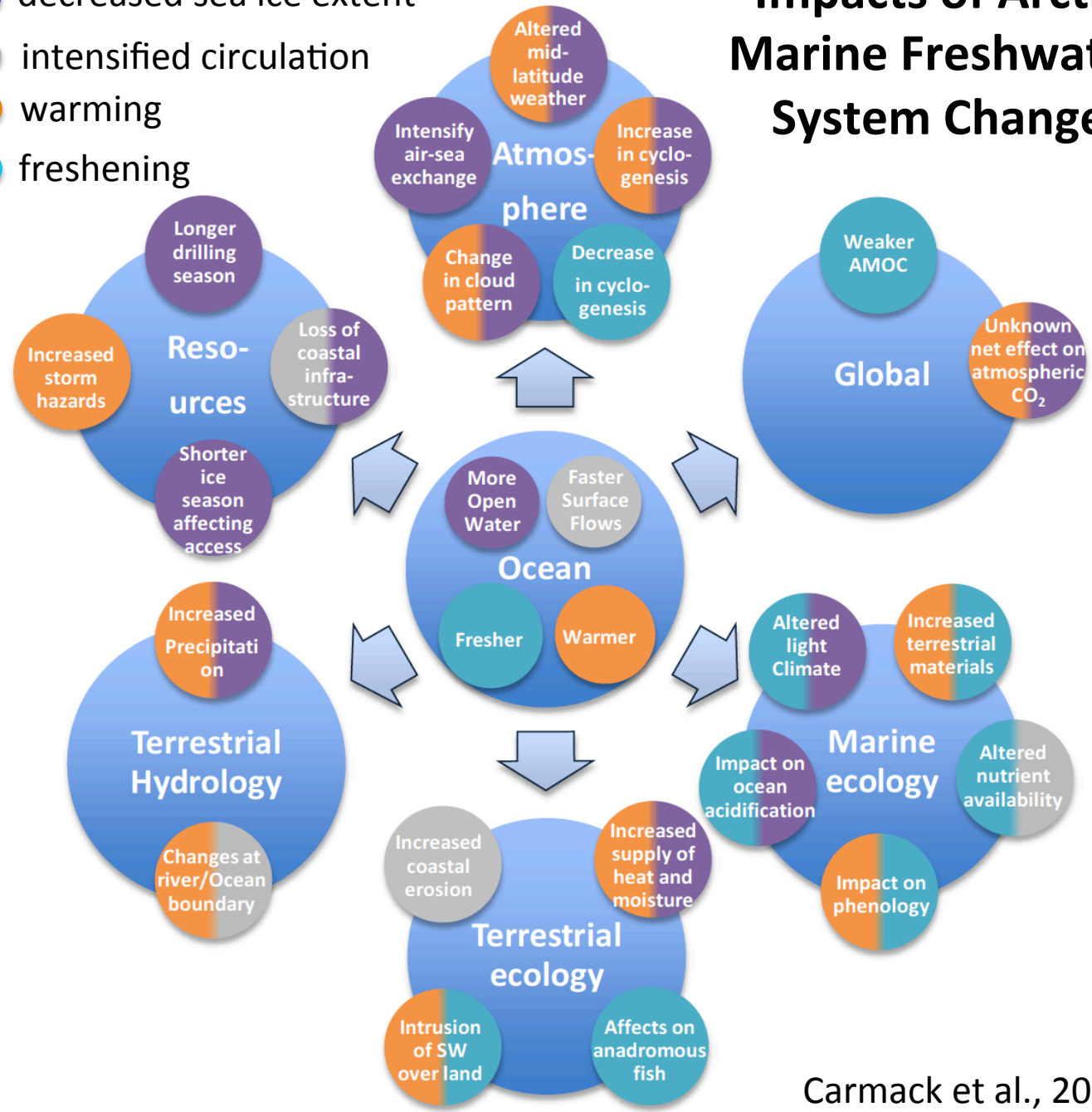
- Salinity controls Arctic Ocean stratification (' β ocean', Carmack, 2007).
- Cold, buoyant surface layer isolates sea ice from heat stored below.
- Modulates sea ice evolution, coupling between atmosphere & ocean.

Global Impacts

- Sea ice growth and melt important component of meridional atmospheric energy transport (Nakamura and Oort, 1988).
- Atlantic Meridional Overturning Circulation (AMOC) sensitive to buoyant Arctic outflow into deepwater formation regions (e.g. Holland et al., 2001; Arzel et al., 2008).
- Changes in Arctic freshwater outflow also modulate:
 - Extent and strength of the North Atlantic subpolar gyre.
 - Northward penetration of warm subtropical waters. Impacts fisheries (Hátún et al., 2009) and carbon uptake and storage (e.g. Schuster and Watson, 2007).

Impacts of Arctic Marine Freshwater System Change

- decreased sea ice extent
- intensified circulation
- warming
- freshening



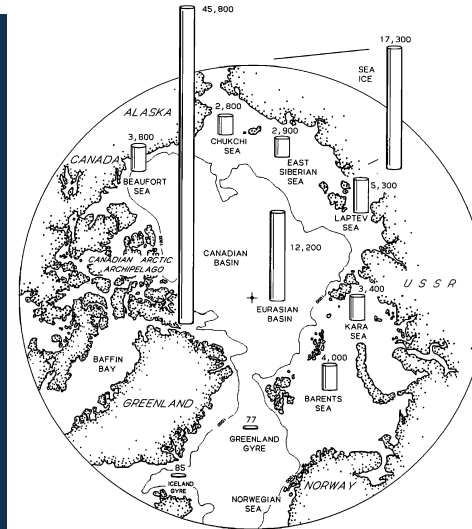
A Freshwater Synthesis: Aagaard and Carmack (1989)

$S_0 = 34.8$
Average Arctic
Ocean salinity

TABLE 1. Fresh Water Budget for the Arctic Ocean

Source or Sink	Transport, $\text{km}^3 \text{ yr}^{-1}$	Yield, cm yr^{-1}
Ice export through Fram Strait	-2790	-29
Water export through Fram Strait	-820	-9
Runoff	3300	35
Precipitation less evaporation	900	9
Water import through Bering Strait	1670	18
Water export through Canadian archipelago	-920	-10
Import with Norwegian Coastal Current	250	3
Saline water import through Barents Sea	-540	-6
Saline water import with West Spitsbergen Current	-160	-2
Net	890	9

- Sparse data – surveys (snapshots), some time series.
- Measurements not contemporaneous.
- Large uncertainties, difficult to discern change.

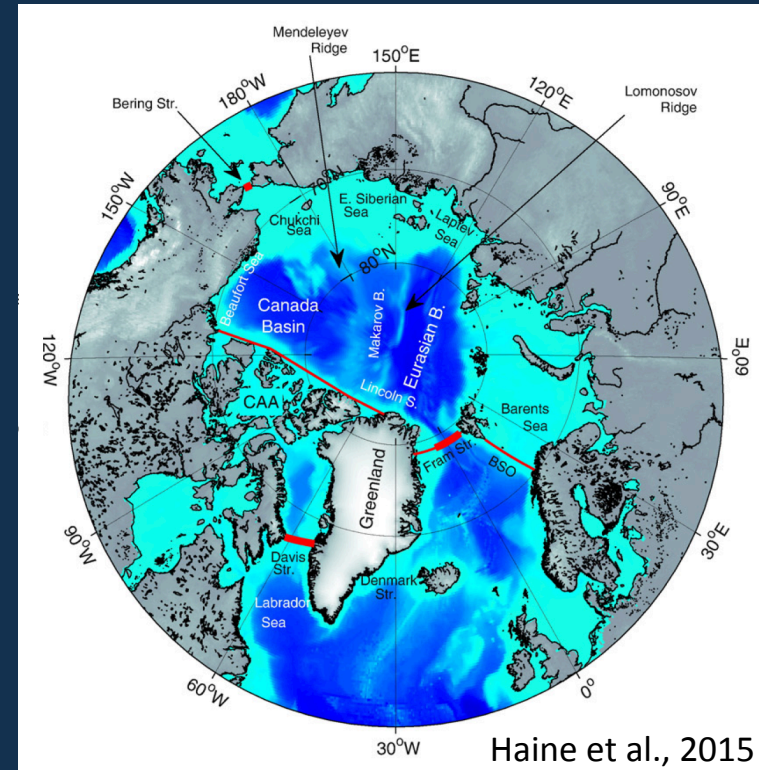


$$\text{FW}_{\text{transport}} = \sum V_i \frac{S_0 - S_i}{S_0} A_i$$

Storage (liquid): 80,000 km^3
Storage (ice): 17,300 km^3

Challenges

- Distributed water column measurements in Arctic interior.
- Resolve dynamically wide straits.
- Measure near ice-ocean interface.
- Sea ice volume.
- Resolve broad range of timescales... seasonal (and shorter) to interannual.
- System undergoing rapid change during observing period.
- Measurements entire FW system.
- Constrain uncertainties to resolve anticipated changes.
- Sustained measurements required to resolve secular change.



Serreze et al. 2006 (NSF FWI)

Inflow

- River runoff ($3900 \pm 390 \text{ km}^3/\text{yr}$)
- Bering Strait liquid ($2400 \pm 300 \text{ km}^3/\text{yr}$)
- P-E ($2000 \pm 200 \text{ km}^3/\text{yr}$)
- Greenland melt ($330 \pm 20 \text{ km}^3/\text{yr}$)
- Bering Strait sea ice ($140 \pm 40 \text{ km}^3/\text{yr}$)

Storage

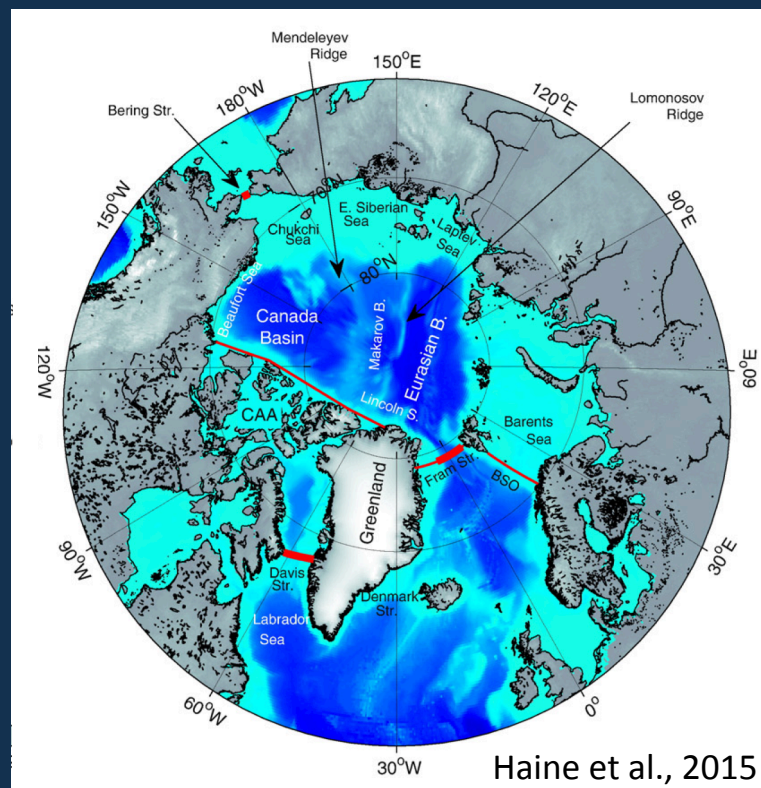
- Liquid FW ($93,000 \text{ km}^3$)
- Seasonal sea ice ($13,000 \text{ km}^3$)
- Multi-year sea ice ($10,900 \text{ km}^3$)

UP
DOWN
NO EST

Outflow

- CAA/Davis Strait liquid ($-3200 \pm 320 \text{ km}^3/\text{yr}$)
- Fram Strait liquid ($-2700 \pm 530 \text{ km}^3/\text{yr}$)
- Fram Strait sea ice ($-2300 \pm 340 \text{ km}^3/\text{yr}$)
- Hudson Strait ($-200 \pm ? \text{ km}^3/\text{yr}$)
- Davis Strait sea ice ($-160 \pm ? \text{ km}^3/\text{yr}$)
- Barents Sea Opening ($-90 \pm 90 \text{ km}^3/\text{yr}$)

- 1980-2004 composite.
- Terrestrial and oceanic measurements with reanalysis products.
- Exploits early data from gateway moorings.
- Interior from PHC climatology.
- Large jump in available data.



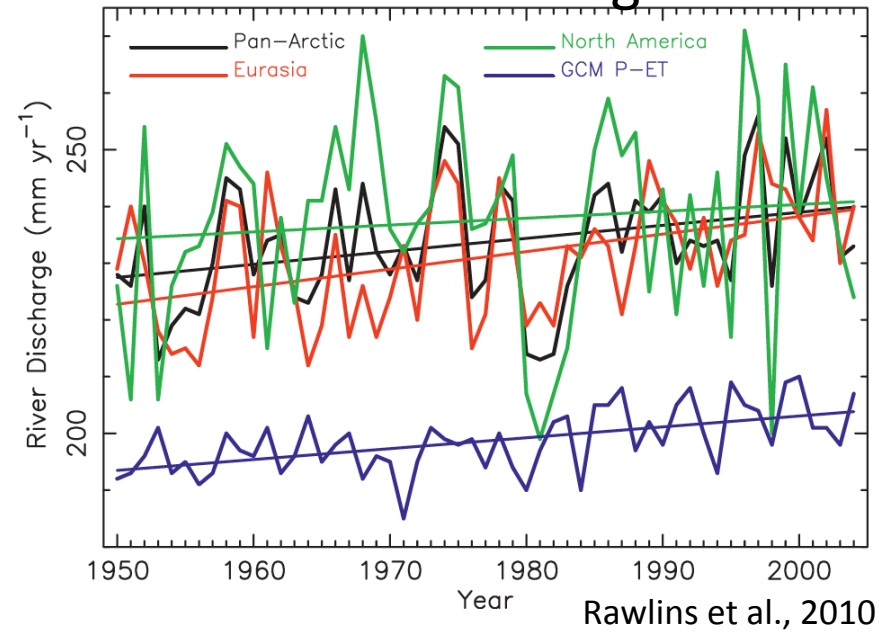
Quantifying River Discharge

Major River Inflows and Pathways



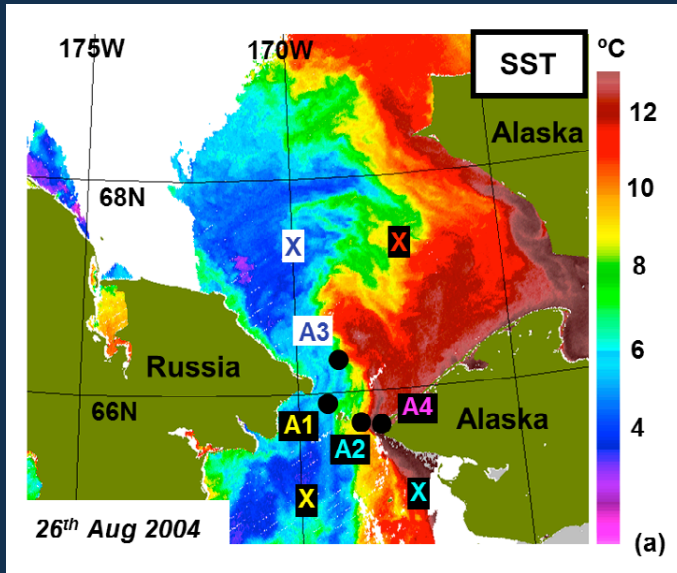
Carmack et al., 2015

River Discharge

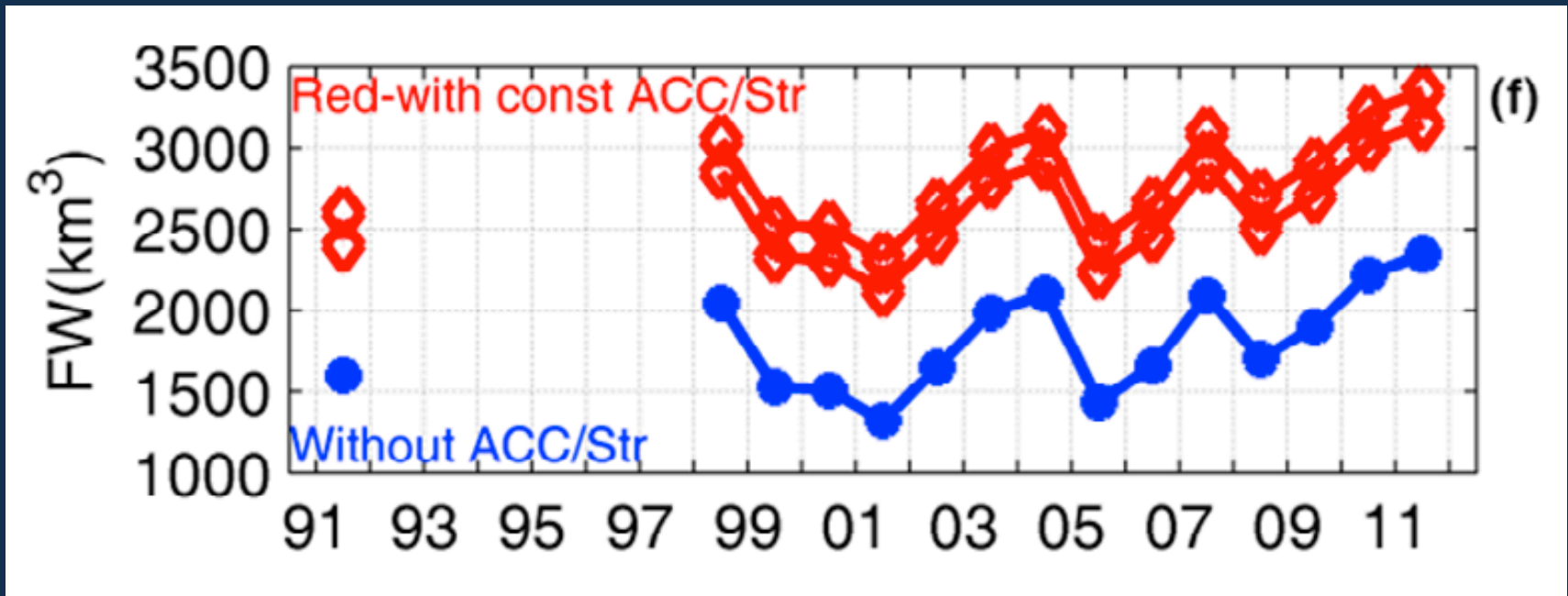


- Major rivers gauged.
- Long records.

Bering Strait



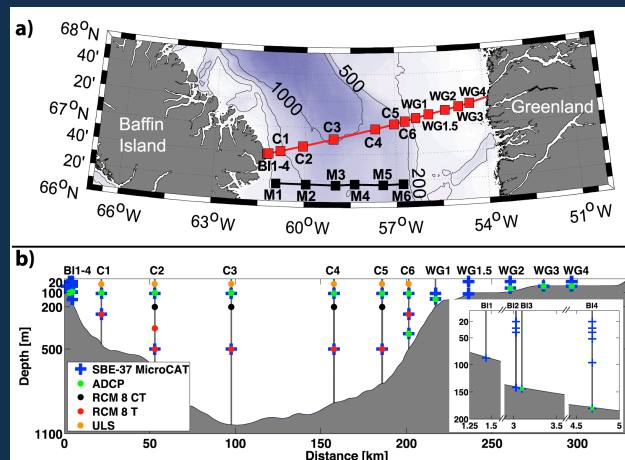
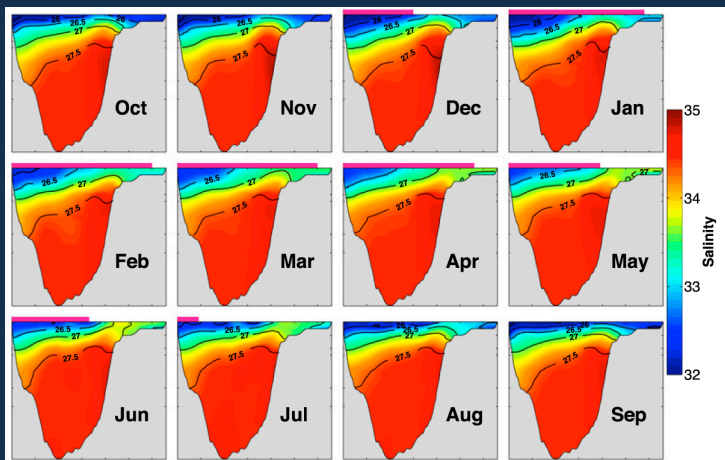
- Corrections for shallow Alaska Coastal Current (ACC).
- ACC resolved in more recent measurements.
- Variability exceeds other inputs.
- Measurements to 2015, future status TBD.



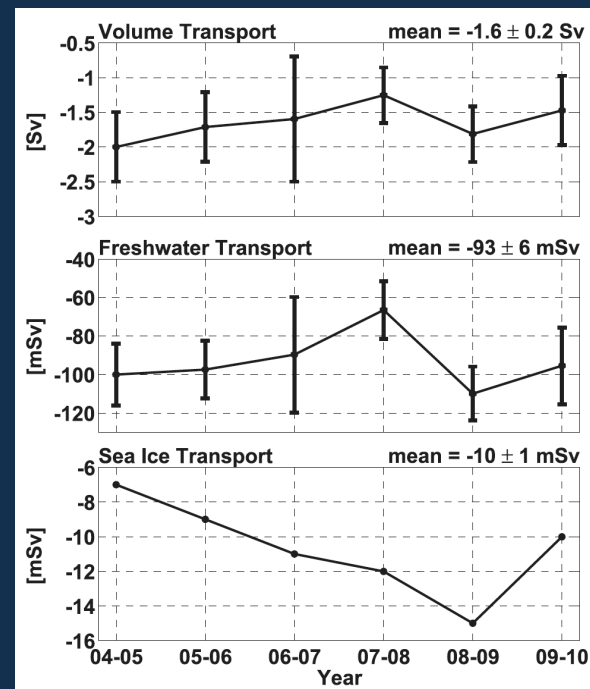
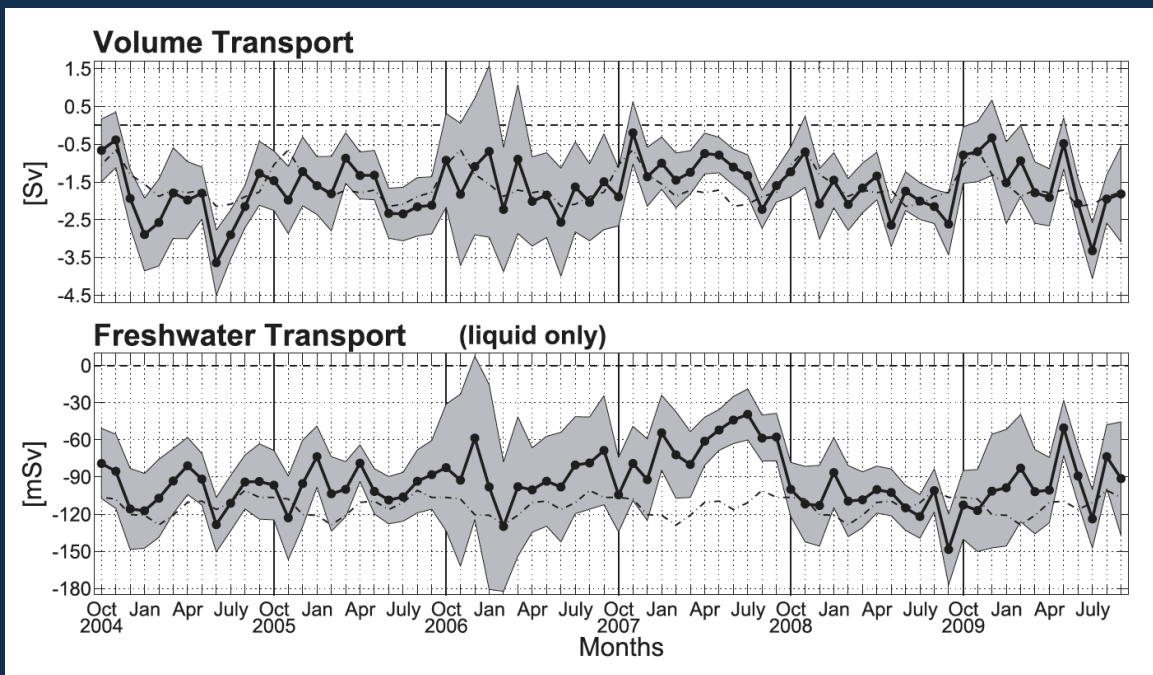
Davis Strait

C. Lee, J. Gobat, B. Curry
APL-UW

B. Petrie, K. Azetsu-Scott
BIO



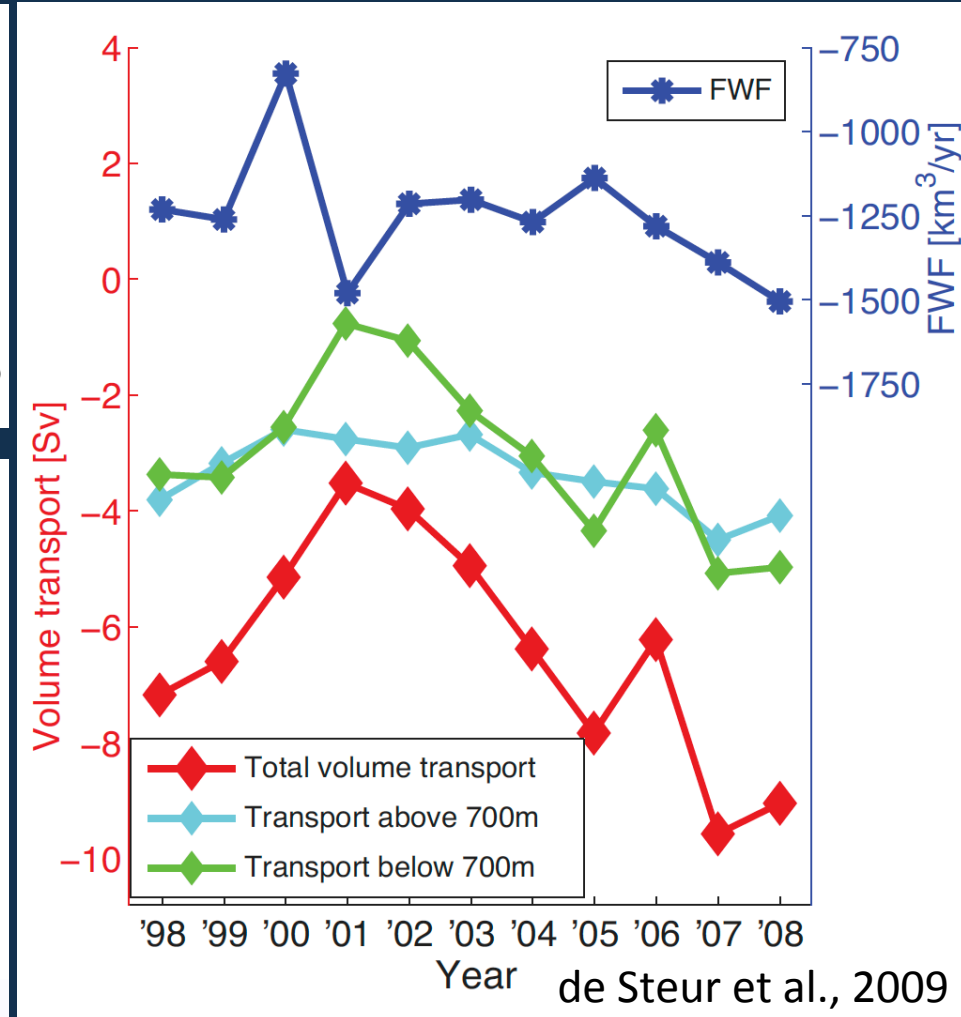
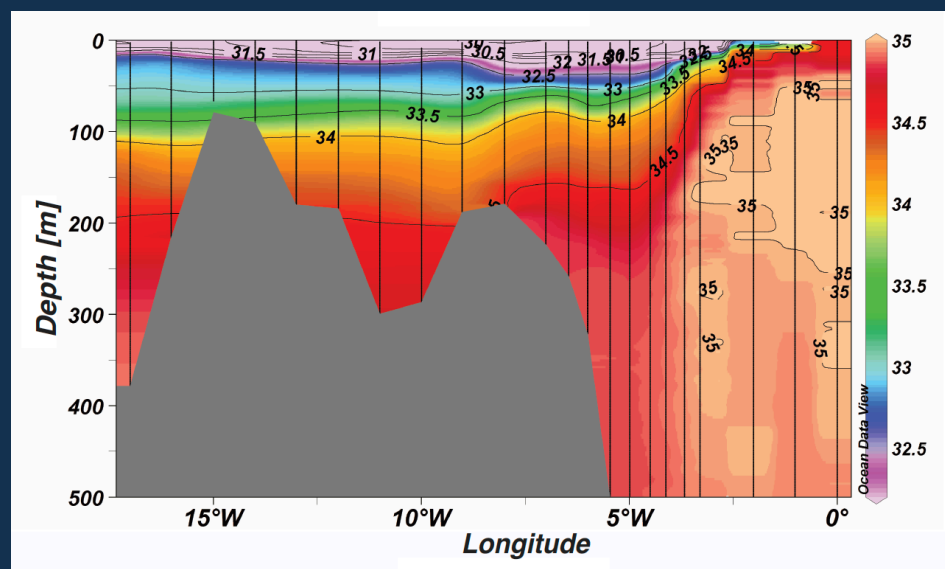
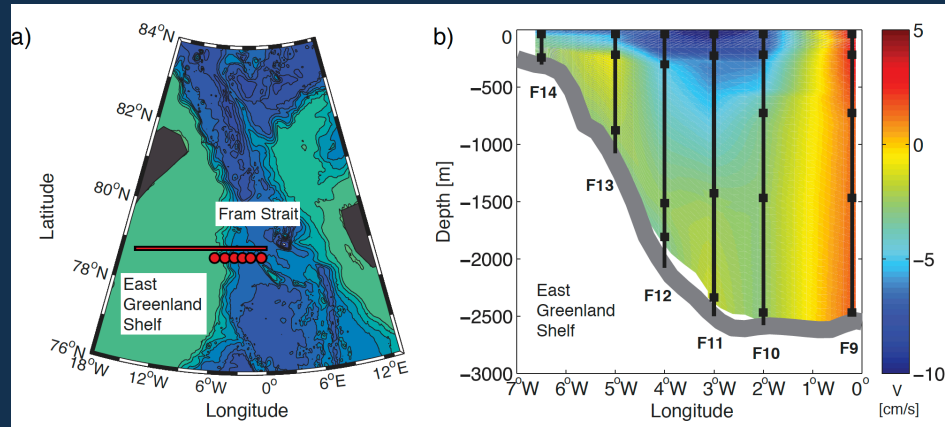
- Mooring, gliders & hydrography.
- 2004 -2015
- Future status TBD



Fram Strait

E. Hansen, L. de Steur, P. Dodd
Norwegian Polar Institute

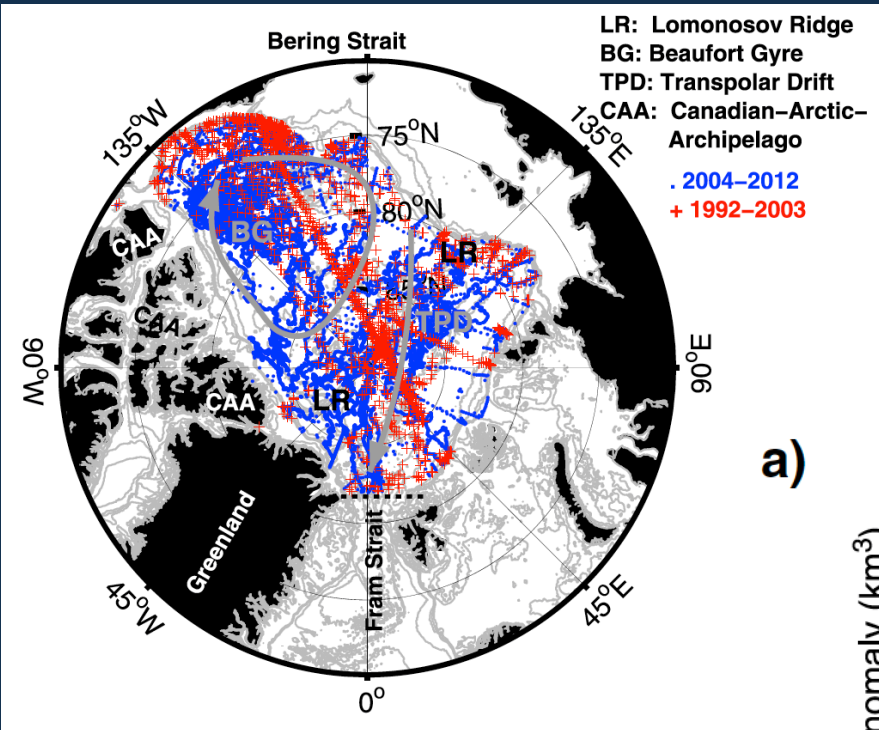
U. Schauer, A. Beszczynska-Moller, T. Kanzow, E. Fahrbach
Alfred Wegener Institute



- East Greenland Shelf, ice-ocean interface poorly resolved.
- Measurement program ongoing as of 2015.

de Steur et al., 2009

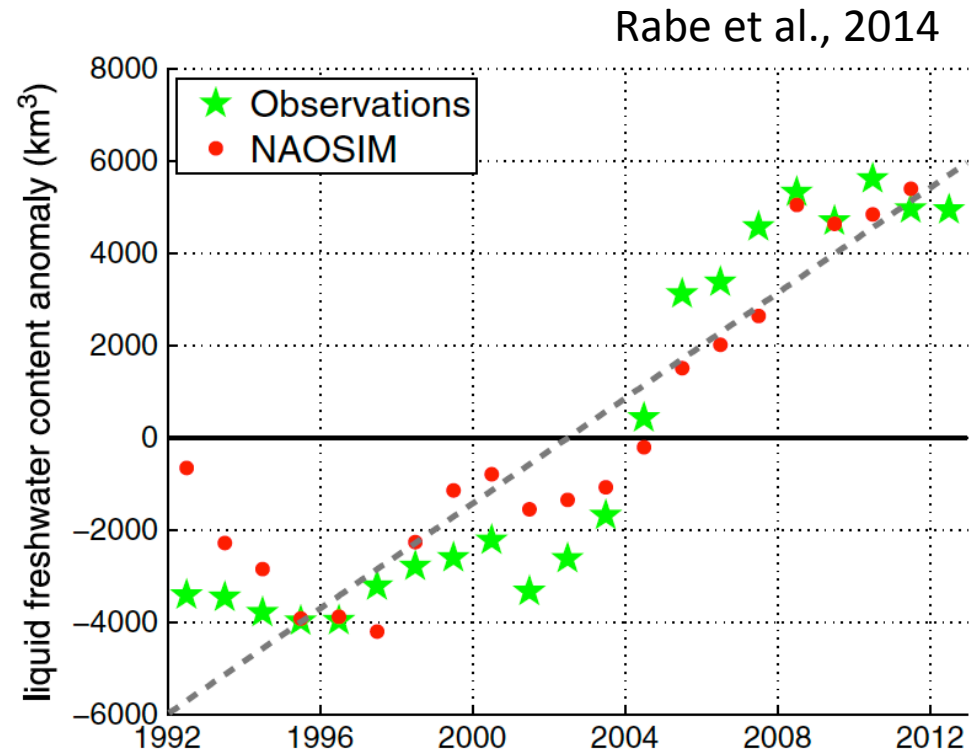
Storage



- Good agreement between **observational estimates** and **NAOSIM model**.
- Trend $600 \pm 300 \text{ km}^3/\text{yr}$.
- Increased Bering inflow, decreased Davis outflow, increased multi-year ice melt?

- Distributed measurements from ships, Ice Tethered Profilers and other buoys.
- Increased autonomous sampling critical.
- $S_0 = 35$

a)

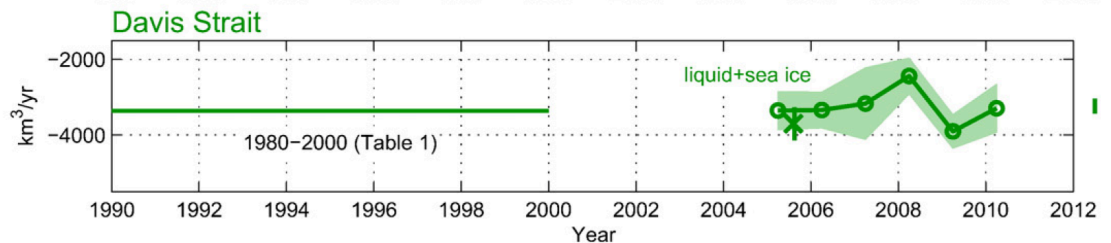
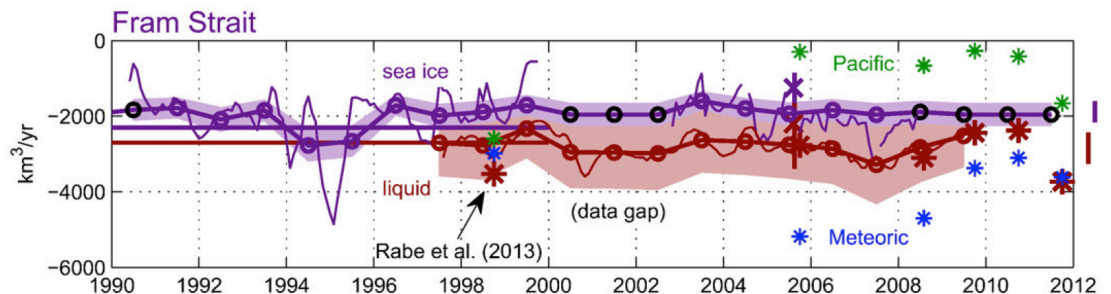
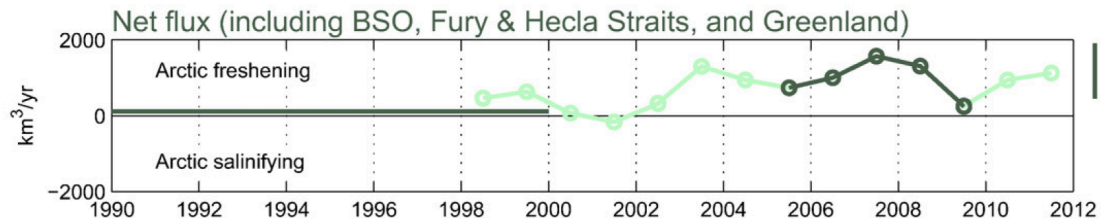
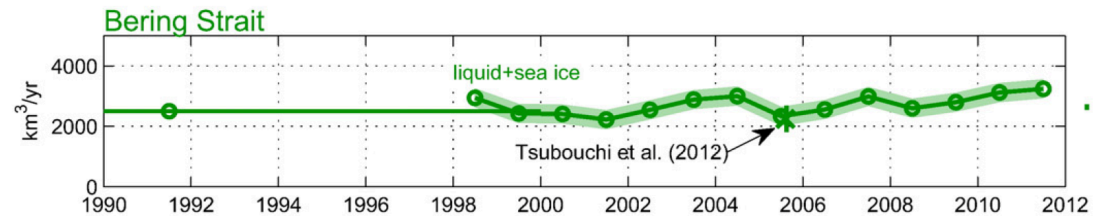
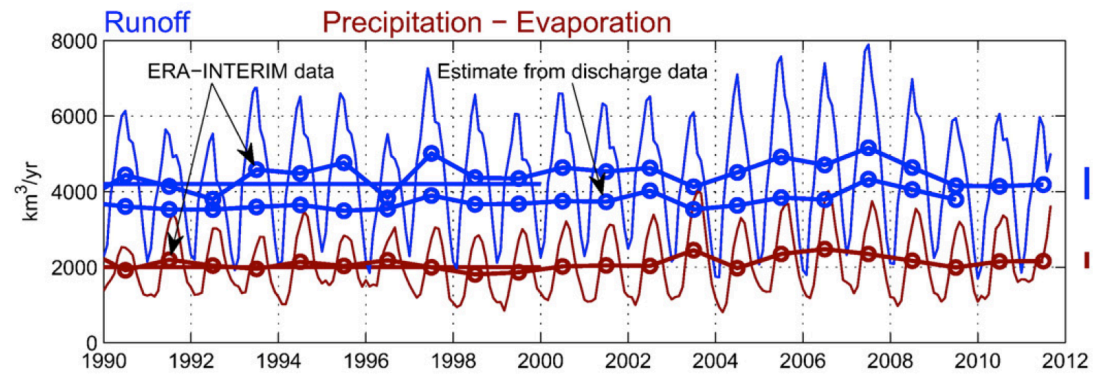
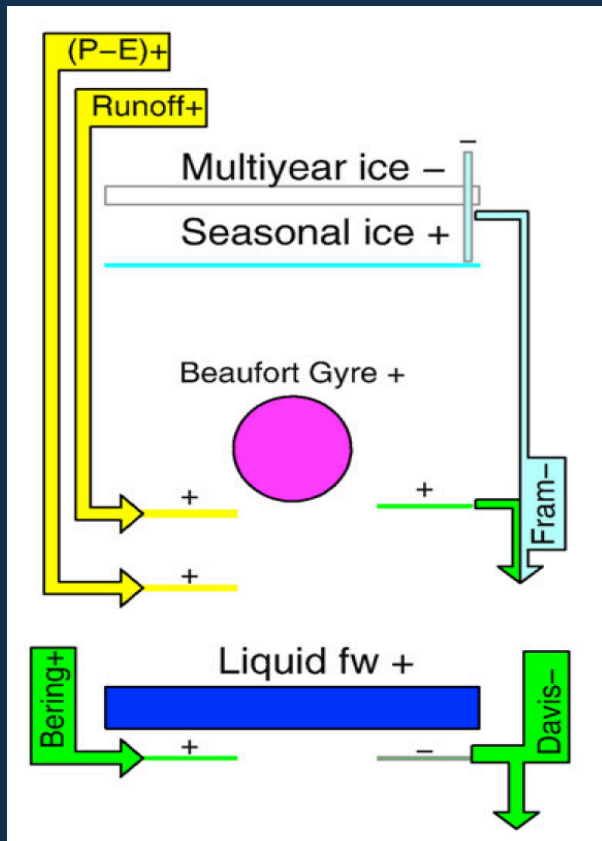


Freshwater Budget

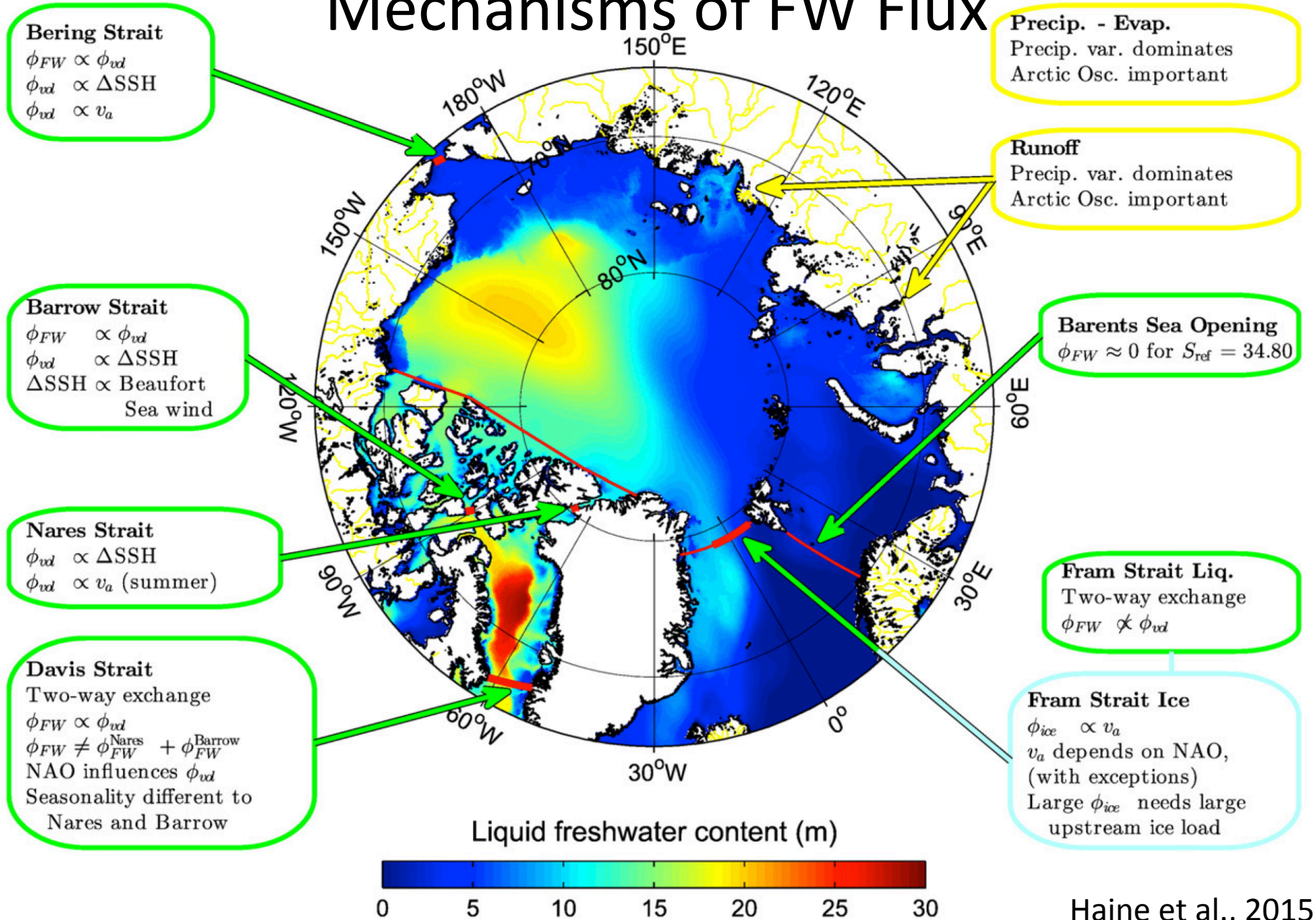
Haine et al., 2015

- 2004-2010 with all gateways quantified.
- Ice storage from PIOMAS assimilation product.

Difference
2000-2010 vs 1980-2000



Mechanisms of FW Flux



Conclusions

FW budgets based on contemporaneous timeseries indicate:

- FW is accumulating in the Arctic, CAA and Baffin Bay.
- Changes in Bering inflow and Fram + Davis outflow below uncertainties.
- Observed increase in storage consistent with increase runoff and P-E, and loss of FW as sea ice.
- Surface winds exert strong controls on FW export and storage.
- No significant change in FW export (Fram + Davis), but release likely to occur in response to changes in wind patterns.

Sustained, contemporaneous measurements of all primary components allows pan-Arctic inverse calculations (e.g. Tsubouchi et al, 2012) to produce self-consistent, pan-Arctic budgets.

Understanding the Arctic FW system requires a network

